

Phytoplankton types and primary production in the Argentine Sea

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ABSTRACT: Several satellite models classify phytoplankton functional types (PFT) based on cell size. In this study we used field data from the Argentine Sea on both the photosynthetic and the bio-optical properties of phytoplankton to distinguish photosynthetic and bio-optical phytoplankton types (PBPT). Cluster analyses were run using data from 70 stations sampled during 3 periods to distinguish different PBPT, and principal component analysis was used to describe them. We examined the main taxonomic composition and percentage of chl *a* in the <5 μm size fraction found within the PBPT. The distribution of PBPT in relation to hourly primary production and environmental conditions was also investigated. The results showed a high degree of variability in bio-optical and photosynthetic properties, e.g. the specific absorption coefficient of phytoplankton, $a_{ph}^B(440)$, varied between 0.015 and 0.067 m^2 (mg chl *a*)⁻¹, and the maximum production at light saturation, P_m^B , varied between 0.68 and 10.05 mg C (mg chl *a*)⁻¹ h⁻¹. This resulted in the discrimination of 11 PBPT. Some had similar average cell sizes but differed in their bio-optical or photosynthetic characteristics, e.g. PBPT1 (with diatoms <5 μm and *Emiliana huxleyi* 2–5 μm) and PBPT6 (with diatoms <5 μm and coccal cells ~2 μm) had markedly different P_m^B values (PBPT1: 1.20 mg C (mg chl *a*)⁻¹ h⁻¹ and PBPT6: 6.71 mg C (mg chl *a*)⁻¹ h⁻¹). This variability in the bio-optical and physiological properties is most likely the result of adaptation by phytoplankton communities to the high heterogeneity in environmental conditions in this region. These results indicate that satellite models describing the distribution of PFT based on cell size alone will not provide a realistic representation of the phytoplankton composition in this highly productive and heterogeneous area.

KEY WORDS: Phytoplankton functional types · Photosynthetic and bio-optical phytoplankton types · Cell size · Primary production · Argentine Sea

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INTRODUCTION

Global biogeochemical as well as trophic web models require information on phytoplankton biomass and on the role of the different groups that compose it in nature. However, obtaining information on phytoplankton diversity using classic taxonomic

methods requires the dedicated work of an expert for an entire day or more for each sample. Therefore, although microscopic analysis remains the most trustworthy way to determine phytoplankton taxonomic composition, several alternative techniques have been developed that classify phytoplankton into groups, commonly known as phyto-

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plankton functional types (PFT). Phytoplankton in a PFT share common ecological roles in nature but do not necessarily have the same phylogenetic origin (Nair et al. 2008). The PFT can be distinguished according to different criteria. Le Quéré et al. (2005) defined functional types of plankton (phyto- and zooplankton) using a combination of different properties, including their biogeochemical role (e.g. N₂ fixers, calcifiers, DMS producers, silicifiers), their nutritional requirement, and the behavior of one group with respect to others.

There are models that differentiate phytoplankton types at large scale (regional and global), from satellite data. Some of these models use algorithms that allow differentiation between individual taxonomic groups, such as coccolithophores (Brown & Podestá 1997), the cyanobacteria *Trichodesmium* spp. (Subramaniam et al. 1999), and diatoms (Sathyendranath et al. 2004), based on their unique optical characteristics. Recently, models have been developed that attempt to describe the distribution of all phytoplankton forms, grouping them into different PFT at both regional and global scales. In these models, the criterion to distinguish phytoplankton types is usually related to cell size (micro-, nano- and picophytoplankton). The principle is based on differentiating, indirectly, cell size through the composition of phytoplankton pigments (by HPLC) and associating these size fractions to some particular bio-optical characteristic that can be derived from satellite estimations, such as radiance, phytoplankton absorption coefficient and chl *a* concentration (Alvain et al. 2005, Uitz et al. 2006, Hirata et al. 2008). Nevertheless, within the same size fraction (e.g. nanoplankton), there could be phytoplankton with a diversity of biogeochemical functions, such as DMS producers and calcifiers (Nair et al. 2008). Some remote sensing studies (e.g. Raitsos et al. 2008) have attempted to distinguish different phytoplankton functional groups within the same size class (for example, diatoms and dinoflagellates). Some of these models have shown a good agreement between satellite and field estimates for some regions such as the southern Benguela ecosystem (Aiken et al. 2007). However, many regions of the world ocean, including the Argentine Sea, have no field data for validation. The importance of obtaining more *in situ* observations to advance these models has been highlighted by many authors (e.g. Brewin et al. 2011).

Photosynthetic and bio-optical properties change according to the composition, abundance and size structure of the phytoplanktonic communities, and with the environmental conditions (e.g. light, tem-

perature, availability of nutrients). For example, values for the photosynthetic parameters α^B (initial slope of production versus irradiance at low values, normalized by chl *a*) and P_m^B (maximum production at light saturation, normalized by chl *a*) usually show spatial and seasonal variations (Platt et al. 1992). An inverse relationship was found between cell size in diatoms and the value of P_m^B by Geider et al. (1986) and Bouman et al. (2005), though this may not always be the case. Several studies have found that small cells (picophytoplankton) have higher values for photosynthetic parameters than large cells (e.g. Platt et al. 1983, Geider et al. 1986). Variations in cell size and pigment composition in different species, as well as in their photophysiological state, influence the shapes of the absorption spectra of phytoplankton. The specific absorption coefficient at wavelength (λ) of phytoplankton (absorption per unit chl *a*), $a_{ph}^B(\lambda)$, is known to decrease with an increase in cell size and intracellular pigment concentration (Sathyendranath et al. 1987). The ratios of absorption at different wavelengths, especially $a_{ph}(440)/a_{ph}(676)$, $a_{ph}(490)/a_{ph}(555)$ and $a_{ph}(443)/a_{ph}(555)$, have been used as an indication of dominant phytoplankton groups (Stuart et al. 2000). The ratio *Chla/FI* (chl *a* concentration versus *in vivo* fluorescence) has also been related to the phytoplankton composition and its physiological state (Kiefer 1973).

The influence of environmental conditions on photosynthetic parameters has been extensively documented, particularly that of the light field (Falkowski 1980), and that of temperature in the case of P_m^B (Bouman et al. 2005), as well as a dependency of α^B on nitrate concentrations (Platt et al. 1992).

The southwestern Atlantic Ocean, and the Argentine Sea in particular, are among the richest chl *a* areas globally according to ocean color satellite images (e.g. Gregg et al. 2005, Rivas et al. 2006). Recent field work has shown wide ranges in primary production during spring (Schloss et al. 2007, Garcia et al. 2008, Lutz et al. 2010). Phytoplankton in the area show a high degree of heterogeneity in their distribution, not only in biomass, but also in community structure (e.g. Negri et al. 1992). Maxima of chl *a* have been associated with different taxonomic groups: diatoms and dinoflagellates in the north (above 41° S; Carreto et al. 1981); diatom blooms in the Argentine shelf break (Lutz & Carreto 1991, Garcia et al. 2008), including one of *Thalassiosira* cf. *oceanica* (Sabatini et al. 2012); the dinoflagellate *Proocentrum minimum* in Bahia Grande ~51° S (Gómez et al. 2011, Sabatini et al. 2012); and *Synechococcus* sp. (picoplankton fraction) during the summer

months at the coastal time-series station EPEA (38° 28' S, 57° 41' W; Silva et al. 2009).

Studies describing phytoplankton according to their function, PFT, are just beginning in the Argentine Sea. Moreno et al. (2012) described PFT in 27 stations in this area during the end of summer/beginning of fall using pigment composition to infer cell size. They found that diatoms dominated the microphytoplankton south of the Argentine shelf break, dinoflagellates occurred close to Peninsula Valdés, nanophytoplankton was abundant in the middle shelf and close to Mar del Plata, and picophytoplankton were found in frontal areas in the south. Using a satellite model, Omachi et al. (2010) showed a wide spatial heterogeneity in PFT prescribed according to cell size (micro-, nano- and picophytoplankton) in different zones of the Argentine Sea from the coast (up to 50 m in depth), middle shelf (up to 200 m) and the shelf break (beyond 200 m).

Most of the studies assigned PFT based on the criterion of cell size, rather than on a given function as defined by Le Quéré et al. (2005). In this study, we considered the following aspects in the distinction of phytoplankton types: (1) that there is an interest in studying the distribution of phytoplankton types using bio-optical remote sensing information, (2) that satellite models of primary production require the assignment of photosynthetic parameters (which can only be estimated by field work), and (3) that a relationship between bio-optical and photosynthetic characteristics could be expected. Thus, our main objective was to distinguish photosynthetic and bio-optical phytoplankton types (PBPT) using bio-optical and photosynthetic field measurements in the Argentine Sea. The main taxonomic composition and the percentage of concentration of chl *a* <5 µm were examined to investigate the predominant cell size within the PBPT. Finally, the relationship between the distribution of these PBPT and primary production and environmental conditions was explored.

MATERIALS AND METHODS

Measurements

Sampling. Field sampling was carried out on board the RV 'Puerto Deseado' covering an extensive area from about 39° to 55° S and 70° to 55° W (Fig. 1) in different seasons: GEF-1, 8 to 28 October 2005 (spring); GEF-2, 10 March to 1 April 2006 (late summer); and GEF-3, 5 to 25 September 2006 (late winter). At all stations, continuous profiles of tempera-

ture and salinity were recorded by a CTD (Sea-Bird 911), and *in vivo* fluorescence was recorded by a Sea-Point fluorometer attached to the CTD.

Seawater samples were collected at the surface and at 2 selected depths for determination of the following variables (see Lutz et al. 2010 for more details).

Chl *a* concentration. Total (*Chla*) and <5 µm fraction (*Chla*_{<5 µm}) concentrations were analyzed using the fluorometric method of Holm-Hansen et al. (1965) with modifications. The continuous *Chla* profile was estimated from the fluorescence profile and discrete measurements of *Chla* (Lutz et al. 2010). Then, the integrated *Chla* within the euphotic zone and in the whole water column were computed (see list of symbols in Table 1). For GEF-2, a linear fitting to the discrete *Chla* values was used to obtain the continuous *Chla* profile since the *in situ* fluorometer was not available.

Particulate absorption coefficients. Total ($a_t(\lambda)$) and detritus ($a_d(\lambda)$) absorption spectra were obtained following the method of Mitchell (1990), using the coefficients of Hoepffner & Sathyendranath (1992), to account for the pathlength amplification factor. The phytoplankton absorption coefficient was calculated by subtraction: ($a_{ph}(\lambda) = a_t(\lambda) - a_d(\lambda)$). The specific absorption coefficient of phytoplankton was computed, normalizing by *Chla* ($a_{ph}^B(\lambda)$). Selected ratios between phytoplankton absorption at a pair of wavelengths were also computed (Table 1).

Field primary production. Surface seawater samples were inoculated with a solution of sodium bicarbonate (¹³C) to a final enrichment of ~8%, transferred into 500 ml square polycarbonate Nalgene bottles and incubated for 3 h in an incubation box with halogen lamps and water circulation regulated to a temperature close to that of the sea at the point where the samples were taken. The intensity of the light in the bottles (measured using a scalar photosynthetically active radiation [PAR] radiometer, Model QSL2100, Biospherical Instruments) was chosen to get an adequate range for the curve of production versus irradiance (*P/E*). After incubation, the samples were filtered onto pre-combusted Whatman GF/F glass fiber filters, fumed with HCl, dried, and encapsulated in tin capsules. The percentage of ¹³C and the amount of particulate organic carbon were analyzed in a mass spectrometer with isotope ratio mass spectrometry (more details in Lutz et al. 2010). The production at each irradiance was calculated based on the equations of Hama et al. (1983), modified by Fernández et al. (2005). The exponential equation of Platt et al. (1980) was used to fit the *P/E* curve and to calculate

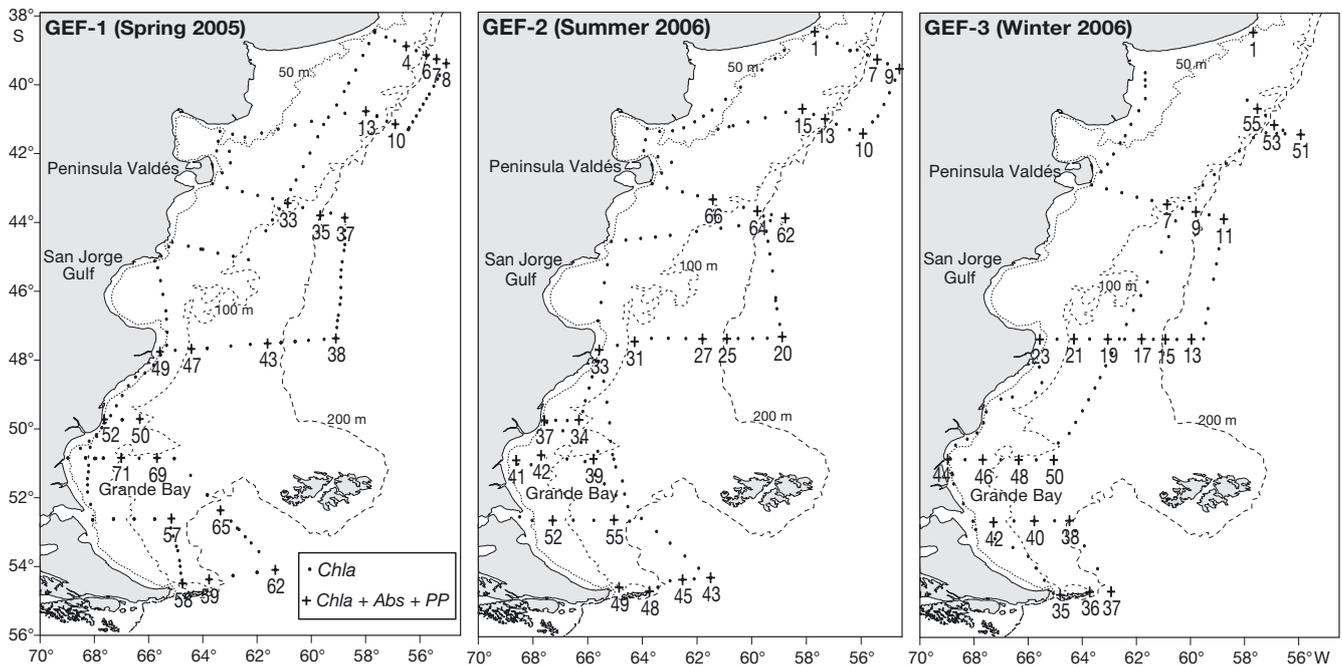


Fig. 1. Location of the stations in the 3 cruises (GEF-1, GEF-2 and GEF-3). Symbols indicate collection of samples for main analyses at the stations. *Chla* represents points where only surface chl *a* was measured, and *Chla* + *Abs* (Absorption) + *PP* (field primary production) represent the primary production stations

the photosynthetic parameters α and P_m . These parameters were normalized by *Chla* to obtain α^B ($\text{mg C (mg chl } a)^{-1} \text{ h}^{-1} (\text{W m}^{-2})^{-1}$) and P_m^B ($\text{mg C (mg chl } a)^{-1} \text{ h}^{-1}$). Hourly production at the surface, p_0 , and daily water-column-integrated primary production, $P_{Z,T}$, were calculated at each station; for the latter, it was assumed that the photosynthetic parameters would remain constant with depth and throughout the day. There are some caveats to these assumptions, e.g. it is known that photosynthetic parameters may change at different depths according to the mixing regime (influencing species distribution and the physiological state of the cells because of changes in light and nutrients; e.g. Platt & Gallegos 1980).

Identification and quantification of surface phytoplankton. Seawater samples were preserved with different fixing solutions: neutralized formaldehyde (0.4% final concentration) for nanophytoplankton (5 to 20 μm) and ultraphytoplankton (<5 μm), and neutralized Lugol's solution for microphytoplankton (>20 μm) (Thronsen 1978). Identification and quantification were carried out using the sedimentation technique and an inverted microscope (Hasle 1978, Tomas 1997). The identification of taxonomic groups was not complete; the abundance of cyanobacteria was not analyzed (since epifluorescence microscopy data were not available), and for GEF-3, only partial information on microphytoplankton was available.

Other environmental measurements. Irradiance at the surface, E_0 , was measured continuously with a LICOR cosine detector. Noon irradiance at the surface for each station, E_n , was estimated as the average of all E_0 recorded within 2 h around local noon and every 2° latitude (Lutz et al. 2010). The downwelling attenuation coefficient of light, $K_d(\text{PAR})$, was computed using the model of Sathyendranath & Platt (1988). The inputs for the model were the absorption and scattering spectra of pure seawater (Pope & Fry 1997), the values of *Chla* and absorption coefficients obtained for discrete depths, and the corresponding average E_0 . Using these values of $K_d(\text{PAR})$ per meter, the irradiance profiles at different times were calculated. The depth of the euphotic zone, Z_{eu} , was taken as the depth at which irradiance, E_z , was equal to 1% E_0 . Wind velocity, W_i , was measured with an on-board meteorological station. The depth of the mixed layer, MLD , was taken as the depth where the difference between the density and the surface density was $>0.05 \text{ kg m}^{-3}$ (Brainerd & Gregg 1995).

Data analysis

Distinction of PBPT. A total of 70 stations with a complete bio-optical and photosynthetic dataset (Fig. 1) from the 3 cruises (22 from GEF-1, 25 from

Table 1. Symbols used, with descriptions and units

Notation	Description	Unit
$a_d(440)$	Absorption coefficient of detritus at 440 nm	m^{-1}
$a_{ph}^B(\lambda)$	Specific absorption coefficient of phytoplankton (absorption per unit chl <i>a</i>) at wavelength (λ)	$\text{m}^2 (\text{mg chl } a)^{-1}$
$\overline{a_{ph}^B}$	Spectrally averaged specific absorption coefficient of phytoplankton	$\text{m}^2 (\text{mg chl } a)^{-1}$
$a_{ph}(440)/a_{ph}(676)$	Ratio between phytoplankton absorption at 440 and 676 nm	Dimensionless
$a_{ph}(490)/a_{ph}(555)$	Ratio between phytoplankton absorption at 490 and 555 nm	Dimensionless
$a_{ph}(443)/a_{ph}(555)$	Ratio between phytoplankton absorption at 443 and 555 nm	Dimensionless
α^B	Initial slope of production versus irradiance (at low values) normalized by chl <i>a</i>	$\text{mg C } (\text{mg chl } a)^{-1} \text{ h}^{-1} (\text{W m}^{-2})^{-1}$
<i>Chla</i>	Total chl <i>a</i> concentration	mg m^{-3}
<i>Chla</i> _{<5μm}	Chl <i>a</i> concentration corresponding to phytoplankton fraction size less than 5 μm	mg m^{-3}
<i>Chla/Fl</i>	Ratio between chl <i>a</i> concentration and <i>in vivo</i> fluorescence	$\text{mg m}^{-3}/\text{relative fluorescence units}$
<i>Chla</i> _{Ze_u}	Chl <i>a</i> integrated to the depth of 1% E_0	mg m^{-2}
<i>Chla</i> _Z	Chl <i>a</i> integrated in the whole water column	mg m^{-2}
% <i>Chla</i> < 5 μm	Percentage of chl <i>a</i> corresponding to phytoplankton fraction size less than 5 μm	%
E_0	Irradiance at surface	$\mu\text{mol quanta m}^{-2} \text{ s}^{-1}$
E_n	Average of E_0 from all measurements recorded every 2° latitude and every 2 h at local noon	$\mu\text{mol quanta m}^{-2} \text{ s}^{-1}$
E_Z	Irradiance at depth	$\mu\text{mol quanta m}^{-2} \text{ s}^{-1}$
$K_d(\text{PAR})$	Diffuse attenuation coefficient for downwelling irradiance in the photosynthetically active range	m^{-1}
<i>MLD</i>	Mixed-layer depth	m
PBPT	Photosynthetic and bio-optical phytoplankton types	
p_0	Hourly surface primary production	$\text{mg C m}^{-3} \text{ h}^{-1}$
P_m^B	Maximum production at saturating irradiance normalized by chl <i>a</i>	$\text{mg C } (\text{mg chl } a)^{-1} \text{ h}^{-1}$
$P_{Z,T}$	Daily water column primary production	$\text{mg C m}^{-2} \text{ d}^{-1}$
<i>Sal</i>	Surface seawater salinity	Practical salinity units
<i>SST</i>	Surface seawater temperature	°C
ϕ_{max}	Maximum quantum yield of photosynthesis	$\text{mol C } (\text{mol quanta})^{-1}$
<i>Wi</i>	Wind velocity	m s^{-1}
<i>Z</i>	Station bottom depth	m
Z_{eu}	Depth at which irradiance, E_Z was equal to 1% E_0	m

GEF-2 and 23 from GEF-3) were used to distinguish PBPT. The photosynthetic parameters used were α^B , P_m^B and maximum quantum yield (ϕ_{max}); and within the bio-optical variables, we chose the following: spectrally averaged specific absorption coefficient ($\overline{a_{ph}^B}$), $a_{ph}^B(440)$, $a_{ph}^B(676)$, $a_{ph}(440)/a_{ph}(676)$, $a_{ph}(490)/a_{ph}(555)$, $a_{ph}(443)/a_{ph}(555)$, total *Chla*, and the ratio *Chla/Fl*. A matrix with the selected standardized properties from all stations in the 3 cruises was built, and a cluster analysis following Ward's method as a criterion to specify the internally most homogeneous group and the square Euclidean distance as a measure of dissimilitude were run using the STATISTICA 7 software package. When the distance of association among stations was <10% of the total cluster distance, the group formed was considered a PBPT.

Characterization of PBPT. First, principal component analysis (PCA) (STATISTICA 7) was applied to all the photosynthetic parameters and bio-optical variables to visualize the degree of similarity between stations forming the PBPT, as well as the extent of correlation between the different properties (Table 2). Next, PBPT were characterized using other phytoplankton properties not used in the distinction of the PBPT such as hourly surface primary production (p_0) and the shapes of the spectra of the specific absorption coefficients $a_{ph}^B(\lambda)$. The relative abundance of each taxonomic group and the percentage of chl *a* < 5 μm (%*Chla*_{<5 μm}) in each sample were used as ancillary information to characterize the PBPT already distinguished by the photosynthetic and bio-optical properties.

Table 2. Ranges of variation, for the 3 cruises, in the photosynthetic and bio-optical properties used in the analyses to distinguish photosynthetic and bio-optical phytoplankton types (except for those marked with †, used to describe the study area)

Variable/ parameter	Range of values
Photosynthetic	
α^B	0.03–0.84 mg C (mg chl a) ⁻¹ h ⁻¹ (W m ⁻²) ⁻¹
P_m^B	0.68–10.05 mg C (mg chl a) ⁻¹ h ⁻¹
Φ_{max}	0.001–0.042 mol C (mol quanta) ⁻¹
† P_0	0.37–55.40 mg C m ⁻³ h ⁻¹
† $P_{Z,T}$	52–5477 mg C m ⁻² d ⁻¹
Bio-optical	
<i>Chla</i>	0.41–19.04 mg m ⁻³
$\overline{Chla}/F1$	0.27–8.79 mg m ⁻³ /relative fluorescence units
$\overline{a_{ph}^B}$	0.006–0.024 m ² (mg chl a) ⁻¹
$a_{ph}^B(440)$	0.015–0.067 m ² (mg chl a) ⁻¹
$a_{ph}^B(676)$	0.006–0.025 m ² (mg chl a) ⁻¹
$a_{ph}(440)/a_{ph}(676)$	1.1–5.2 dimensionless
$a_{ph}(490)/a_{ph}(555)$	2.1–13.2 dimensionless
$a_{ph}(443)/a_{ph}(555)$	3.4–21.7 dimensionless
† $Chla_Z$	27–1486 mg m ⁻²
† $Chla_{Zeu}$	12–277 mg m ⁻²

Relationship between PBPT and environmental conditions. The available environmental properties included basic physical (temperature, salinity, wind speed, total depth, depth of the mixed layer, and incident irradiance) and underwater bio-optical (depth of the euphotic layer, detritus absorption, attenuation coefficient) estimations (Table 3). These were used to characterize the PBPT according to environmental conditions using PCA.

Satellite estimation of phytoplankton size classes. The model to determine phytoplankton size classes from satellite ocean color developed by Hirata et al. (2008) was applied to moderate resolution imaging spectroradiometer (MODIS)-Aqua satellite compos-

ite images of $a_{ph}(443)$ using the quasi-analytical algorithm (Lee et al. 2002) for each cruise. MODIS Level 1A and ancillary data from each cruise period were downloaded from the NASA Web site (<http://oceancolor.gsfc.nasa.gov/>), processed using SeaDAS v6.2 (2009 reprocessing) and mapped to a cylindrical equidistant projection with a spatial resolution of 1 min (~2 km). Maps of phytoplankton size classes were obtained using $a_{ph}(443)$ thresholds to detect the dominance of micro- ($a_{ph}(443) > 0.069$ m⁻¹), nano- ($0.023 < a_{ph}(443) \leq 0.069$ m⁻¹) and picophytoplankton ($a_{ph}(443) \leq 0.023$ m⁻¹).

RESULTS

Chl a and primary production distribution

Chl a concentrations showed a high degree of variation in the 3 periods analyzed (Fig. 2A), with maximum values recorded in spring (24.01 mg m⁻³ at Stn 73, GEF-1, in Grande Bay, with no primary production data) and minimum values recorded at the end of winter (0.42 mg m⁻³ at Stn 51, GEF-3). Marked blooms were observed in spring, while at the end of summer, these were of minor intensity (Fig. 2A). At the end of winter, *Chla* values were low in the south and high in the north, indicating the beginning of spring (GEF-3, Fig. 2A). The concentration of chl a in the <5 μm fraction showed a widespread distribution (according to % *Chla*_{<5μm}), with some areas denoting the predominance of large cells in GEF-1 and GEF-2. At the end of winter (GEF-3), % *Chla*_{<5μm} was above 60% in the region south of 47° S, while in the north, total *Chla* was dominated by the >5 μm fraction (Fig. 2B). The approximated values of integrated production, $P_{Z,T}$, were also highly variable in the 3 cruises (Table 2 & Fig. 2C), with the highest value found in spring ($P_{Z,T} = 5477$ mg C m⁻² d⁻¹ and *Chla* = 19.05 mg m⁻³ at Stn 7, GEF-1) and the lowest found at the end of summer ($P_{Z,T} = 52$ mg C m⁻² d⁻¹ and *Chla* = 0.81 mg m⁻³ at Stn 45, GEF-2). Although there was a significant positive Spearman correlation, r_s , between $P_{Z,T}$ and *Chla* ($r_s = 0.59$, $p < 0.05$) showing some stations with high production and *Chla* values (e.g. Stn 7, GEF-1), there was a marked dispersion of the data around the correlation line (Fig. 3), e.g. Stns 50 and 52 from GEF-1 showed high values of $P_{Z,T}$ (2340 mg C m⁻² d⁻¹ and 1642 mg C m⁻² d⁻¹, respectively), while their *Chla* values were relatively low (1.27 mg m⁻³ and 1.12 mg m⁻³, respectively). Furthermore, for these last stations, *Chla* values were fairly similar, while $P_{Z,T}$ differed by more than 50%.

Table 3. Ranges of the environmental properties in the 3 cruises

Variable/parameter	Range of values
Bio-optical	
Z_{eu}	13–29 m
$K_d(\text{PAR})$	0.163–0.376 m ⁻¹
$a_d(440)$	0.001–0.085 m ⁻¹
Physical	
E_p	157–1365 μmol quanta m ⁻² s ⁻¹
W_i	0.51–14.40 m s ⁻¹
<i>SST</i>	4.388–19.660 °C
<i>Sal</i>	32.304–34.538
<i>Z</i>	23–2980 m
<i>MLD</i>	2–135 m

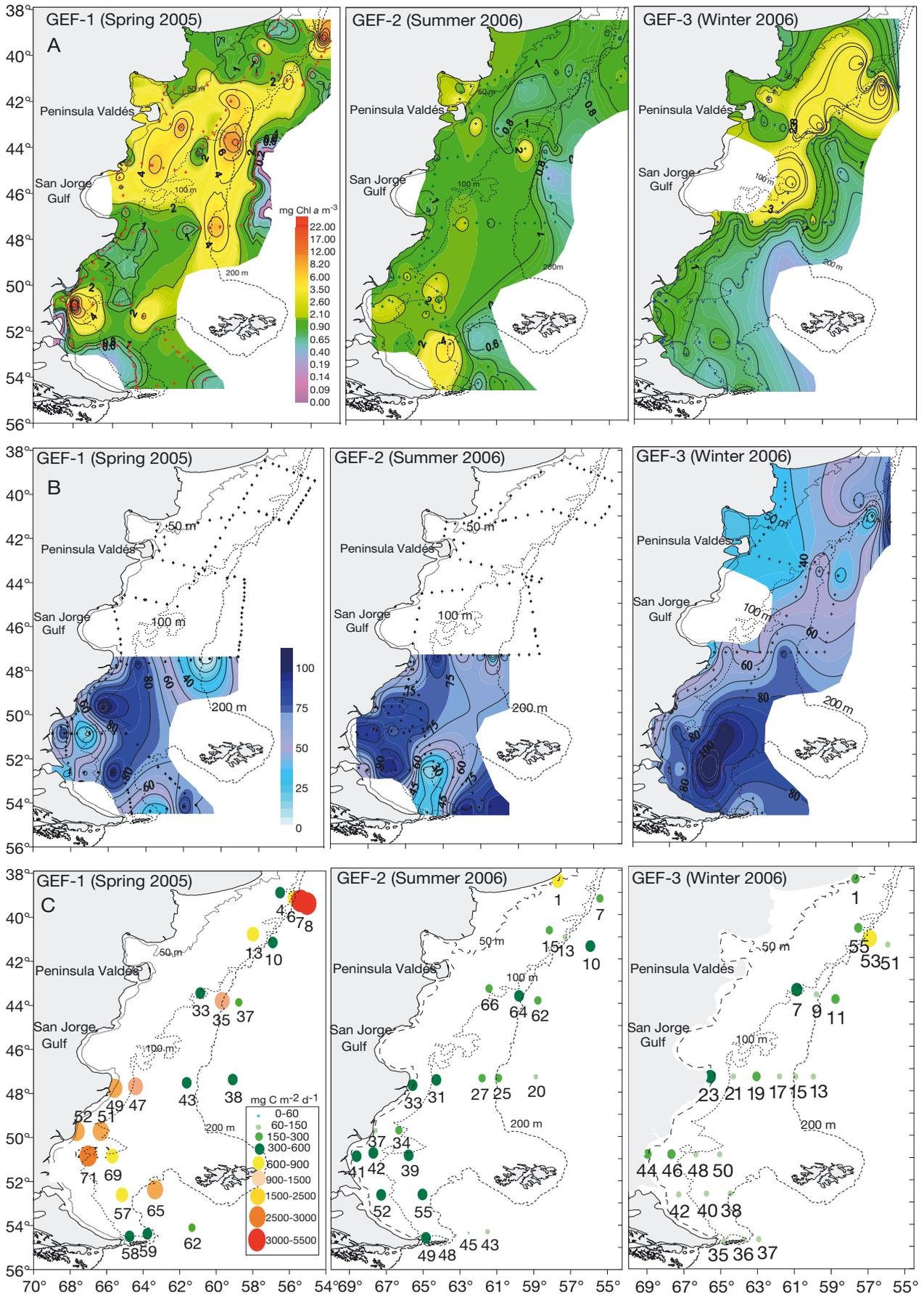


Fig. 2. Distribution of (A) total surface chl a, (B) percentage of surface chl a in the <5 μm fraction, and (C) daily water-column-integrated primary production ($P_{Z,T}$) for the 3 cruises

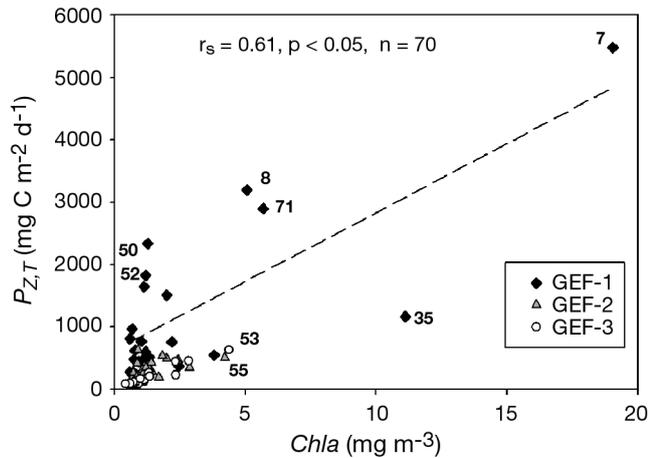


Fig. 3. Spearman correlation between the values of $P_{z,T}$ and $Chla$ for the 3 cruises

Variations in surface bio-optical and photosynthetic properties

High variations in all properties were observed during the 3 cruises (Table 2), e.g. P_m^B varied between 0.68 and 10.05 mg C (mg chl a)⁻¹ h⁻¹, and $a_{ph}^B(440)$ varied between 0.015 and 0.067 m² (mg chl a)⁻¹. Significant correlations were found between $a_{ph}(440)/a_{ph}(676)$ and $Chla$ ($r_s = -0.37$, $p < 0.05$; Fig. 4A), between $a_{ph}^B(440)$ and $Chla_{<5\mu m}$ ($r_s = 0.29$, $p < 0.05$; Fig. 4B) and between $a_{ph}^B(676)$ and $Chla_{<5\mu m}$ ($r_s = 0.52$, $p < 0.05$; Fig. 4C); note that $Chla_{<5\mu m}$ reached values above 2 mg m⁻³ (Fig. 4B,C). However, no significant correlations were found between the photosynthetic parameters (α^B and P_m^B) with respect to total $Chla$ or $Chla_{<5\mu m}$.

PBPT

Distinction and distribution of PBPT. A total of 11 PBPT were distinguished from the cluster analysis

based on the photosynthetic and bio-optical properties for the 3 periods studied (Fig. 5). The spring cruise (GEF-1) showed the highest diversity, 9 PBPT, of which 4 were exclusively from that period. The latter were PBPT2 (diatoms and dinoflagellates—shelf break), PBPT5 (mixed phytoplankton—shelf, shelf break and outer shelf), PBPT6 (ultradiatoms and Chlorophyta—shelf) and PBPT7 (bloom of *Thalassiosira cf. oceanica*—shelf break). GEF-2 and GEF-3 each had 5 PBPT, with only 1 PBPT exclusive for the GEF-3 period, i.e. PBPT11 (ultra- and nanophytoplankton: shelf break) (Table 4 & Fig. 6).

Characterization of the PBPT. Large variations in the average of the photosynthetic parameters, chl a and primary production were found within most of the PBPT formed by 2 or more stations (Table 4). The PCA of the bio-optical and photosynthetic properties showed that 4 principal components (PC) explained 85% of the total variation (Table 5). PC1 was mainly influenced by the specific absorption coefficients, while PC2 was mostly related to the photosynthetic properties (P_m^B , α^B and ϕ_{max}) (Table 5 & Fig. 7A). Finally, PC3 and PC4 were associated with the ratios of absorption between pairs of wavelengths and $Chla$, respectively (Table 5 & Fig. 7C). The projection of the different stations from the different PBPT in the plane of the PCs (Fig. 7B,D) offers a visualization of the association of the PBPT with the different components. For example, PBPT3 and PBPT4 are separated from the rest of the PBPT because of the high values in the specific absorption coefficients (PC1) at the respective stations, in contrast to PBPT8, which has low values (Fig. 7B). Regarding PC2, PBPT5 showed the highest values in photosynthetic parameters and low $Chla/Fl$ (Fig. 7B), while PBPT11 showed extreme values in PC2, with the lowest photosynthetic parameters. Some PBPT, such as PBPT2 and PBPT6, were differentiated from the rest by their association to PC3 (Fig. 7D), showing high values in the ratios of absorption at different wavelengths. The

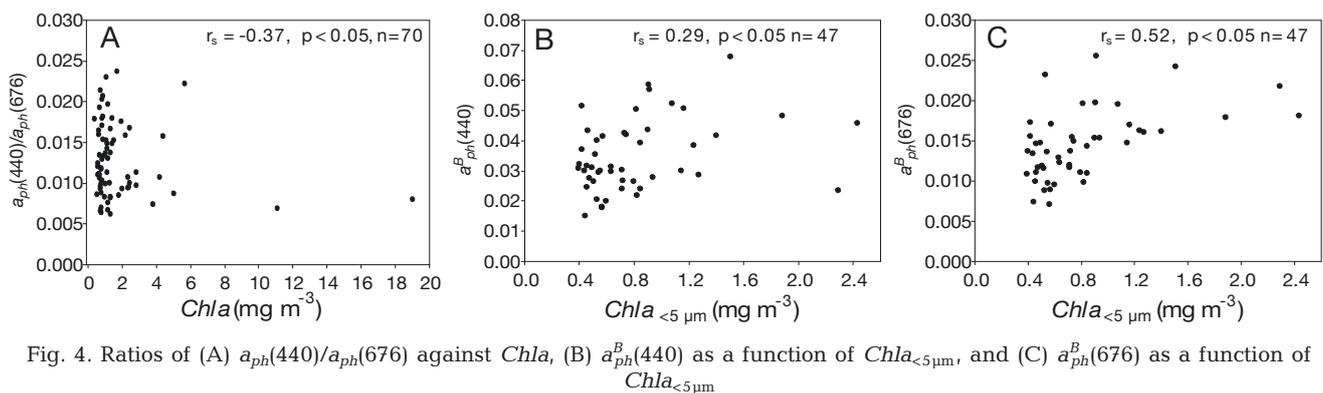


Fig. 4. Ratios of (A) $a_{ph}(440)/a_{ph}(676)$ against $Chla$, (B) $a_{ph}^B(440)$ as a function of $Chla_{<5\mu m}$, and (C) $a_{ph}^B(676)$ as a function of $Chla_{<5\mu m}$

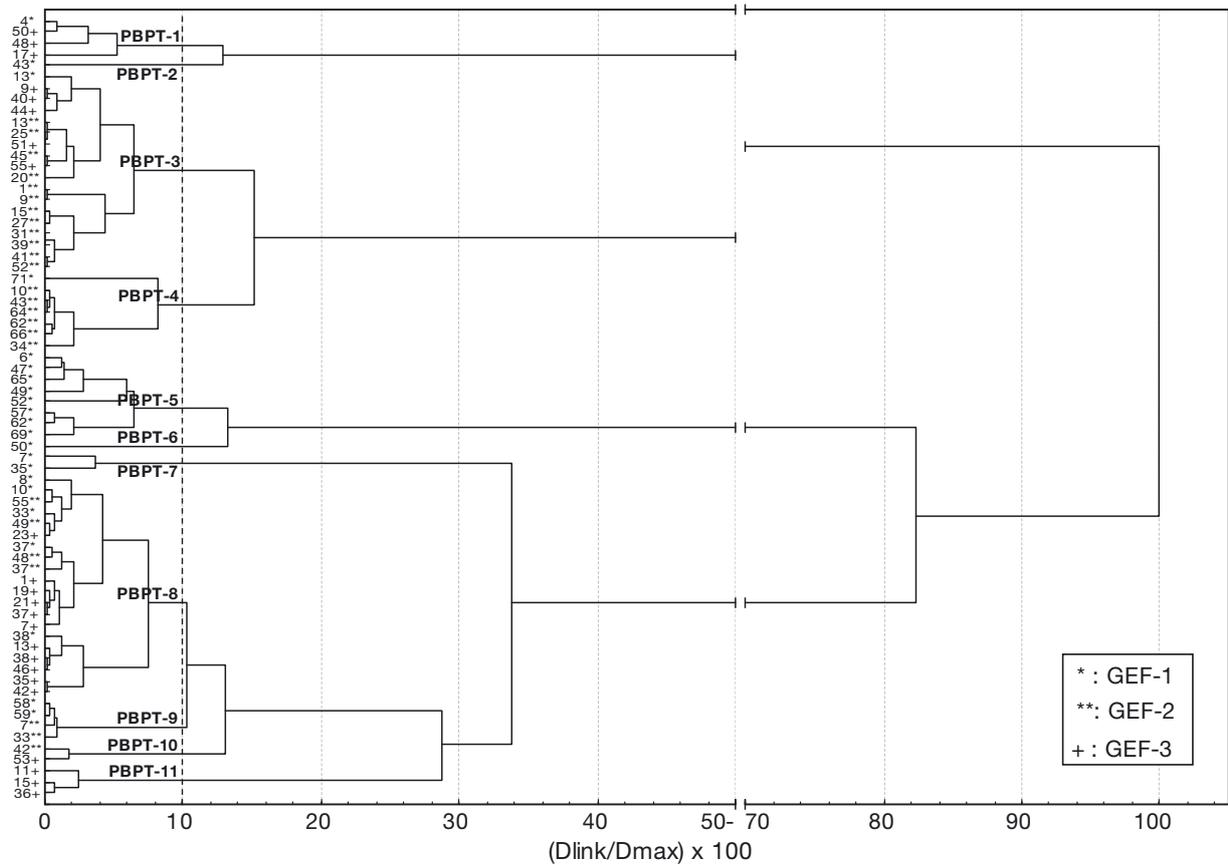


Fig. 5. Cluster analysis to distinguish photosynthetic and bio-optical phytoplankton types (PBPT) based on photosynthetic and bio-optical properties including all stations from the 3 cruises. The dashed line represents the threshold line of the cluster marked at 10% of the total distance to distinguish the PBPT

Table 4. Mean values of the properties analyzed in each photosynthetic and bio-optical phytoplankton type (PBPT). Bold numbers represent properties with coefficients of variation within each PBPT <35%. Note PBPT2 and PBPT6 have only 1 station each

	PBPT1	PBPT2	PBPT3	PBPT4	PBPT5	PBPT6	PBPT7	PBPT8	PBPT9	PBPT10	PBPT11
n (stations):	4	1	18	7	8	1	2	20	4	2	3
Station numbers:	GEF-1: 4	GEF-1: 43	GEF-1: 13	GEF-1: 71	GEF-1: 6, 47, 49, 52, 57, 62, 65, 69	GEF-1: 50	GEF-1: 7, 35	GEF-1: 8, 10, 33, 37, 38	GEF-1: 58, 59	GEF-2: 42	GEF-3: 11, 15, 36
	GEF-3: 17, 48, 50		GEF-2: 1, 9, 13, 15, 20, 25, 27, 31, 39, 41, 45, 52	GEF-2: 10, 34, 43, 62, 64, 66				GEF-2: 37, 48, 49, 55	GEF-2: 7, 33	GEF-3: 53	
			GEF-3: 9, 40, 44, 51, 55					GEF-3: 1, 7, 13, 19, 21, 23, 35, 37, 38, 42, 46			
P_m^B	1.20	1.24	1.87	2.69	6.21	6.71	1.87	1.38	3.86	1.51	0.96
α^B	0.09	0.06	0.07	0.10	0.38	0.84	0.10	0.07	0.09	0.07	0.05
ϕ_{max}	0.004	0.005	0.012	0.005	0.019	0.042	0.005	0.004	0.004	0.004	0.002
p_0	1.16	1.14	2.26	7.24	9.87	8.51	32.31	2.69	4.68	4.34	1.02
<i>Chla</i>	0.83	0.92	1.18	1.72	0.99	1.26	15.00	1.77	1.25	3.38	1.06
<i>Chla/FI</i>	2.00	0.64	1.07	0.71	0.58	0.63	2.37	1.67	1.13	1.08	7.00
a_{ph}^B	0.012	0.013	0.016	0.022	0.011	0.010	0.008	0.010	0.007	0.013	0.010
$a_{ph}^B(440)$	0.039	0.046	0.046	0.061	0.035	0.028	0.019	0.027	0.018	0.032	0.029
$a_{ph}^B(676)$	0.015	0.009	0.016	0.022	0.012	0.016	0.011	0.012	0.008	0.023	0.012
$a_{ph}(440)/a_{ph}(676)$	2.52	5.25	2.91	2.85	2.75	1.78	1.73	2.21	2.34	1.40	2.37
$a_{ph}(490)/a_{ph}(555)$	9.97	8.88	4.16	4.37	4.18	5.62	2.72	4.26	2.72	2.79	5.73
$a_{ph}(443)/a_{ph}(555)$	16.43	17.28	6.33	6.22	6.79	8.53	4.16	6.64	4.19	4.15	9.51

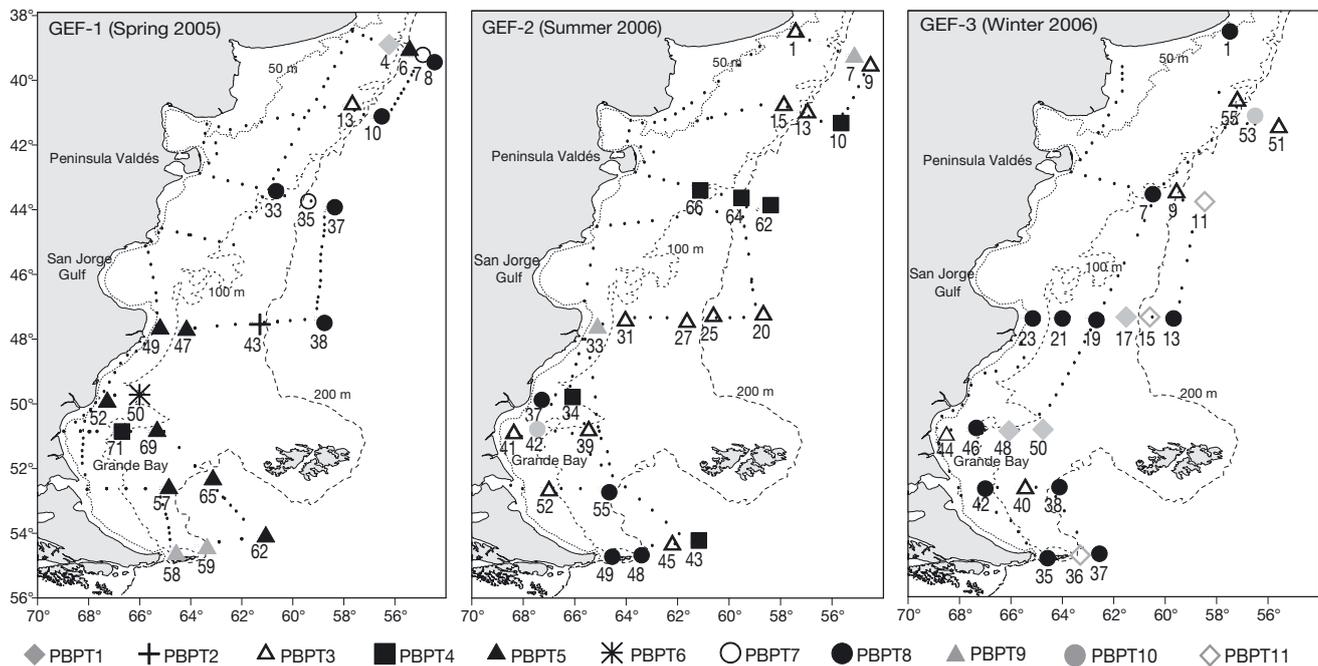


Fig. 6. Distribution of the photosynthetic and bio-optical phytoplankton types (PBPT) in the different periods analyzed

shapes of the absorption spectra at the different stations within a PBPT were usually similar, but some PBPT had strong differences in the blue region, such as in PBPT5 and PBPT10 (Fig. 8).

The available information on taxonomic composition within the PBPT showed a high heterogeneity

Table 5. Pearson correlation coefficients between main principal components (PC) and properties used to separate photosynthetic and bio-optical phytoplankton types. Percentage of variance accounted for by each component is shown in parentheses. Correlation coefficients were significant, above the critical value 0.28, with $p < 0.01$. Highest values of the coefficients in each component are shown in bold

Variable/ parameter	PC1 (30%)	PC2 (25%)	PC3 (20%)	PC4 (10%)
P_m^B	0.42	-0.83		
α^B	0.34	-0.80		
ϕ_{\max}	0.42	-0.83		
$Chla$	0.35		0.40	-0.49
$Chla/Fl$		0.58		-0.29
$\overline{a_{ph}^B}$	-0.87	-0.34	0.28	
$a_{ph}(440)/a_{ph}(676)$	-0.60		-0.30	0.52
$a_{ph}^B(440)$	-0.92	-0.31		
$a_{ph}^B(676)$	-0.69		0.35	-0.52
$a_{ph}(490)/a_{ph}(555)$	-0.35		-0.87	
$a_{ph}(443)/a_{ph}(555)$	-0.30		-0.89	

(Table 6). For example, PBPT1 and PBPT6, with % $Chla_{<5\mu m}$ values above 74 %, both contained high concentrations of diatoms of the ultrafraction (<5 μm), whereas PBPT1 also had *Emiliania huxleyi*, and PBPT6 had coccal cells (Table 6). PBPT7 was composed of a bloom of the diatom *Thalassiosira cf. oceanica* (nanofraction), and PBPT4, with values of % $Chla_{<5\mu m}$ ranging between 16 and 100 %, was formed by 7 stations with different phytoplankton composition: Stn 71 (GEF-1), characterized by a bloom of *Prorocentrum minimum* (nanofraction), and Stns 10, 34, 43, 62, 64, and 66 (GEF-2), principally dominated by Chlorophyta of the ultrafraction (Table 6).

High variability was observed in the values of surface hourly primary production (p_0) within the PBPT. Therefore, no significant correlation was found between mean p_0 and mean $Chla$ for the PBPT. The maximum values of p_0 and of $Chla$ were observed in PBPT7 (spring—bloom of *Thalassiosira cf. oceanica*—shelf break); the minimum value of p_0 was found in PBPT11 (summer—ultra- and nanophytoplankton—shelf break), with $Chla = 1.06 \text{ mg m}^{-3}$; and the minimum $Chla$ value was found in PBPT1 (spring and winter—shelf), with $p_0 = 1.16 \text{ mg C m}^{-3} \text{ h}^{-1}$ (Table 4).

Environmental conditions were highly variable considering the whole area and the 3 periods studied (Table 3). The PCA showed that the first 4 PCs represented ~75 % of the total variation (Table 7). PC1 was

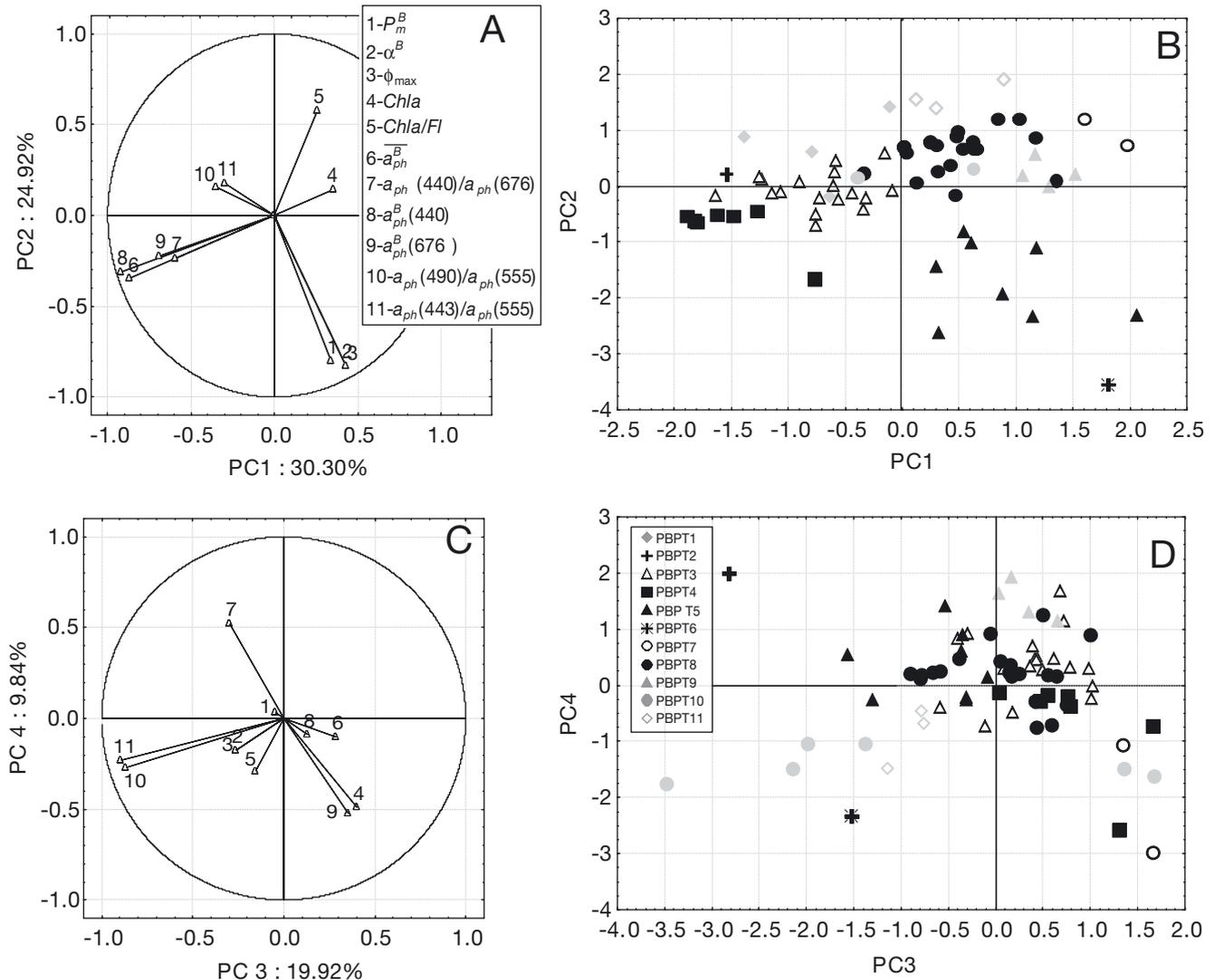


Fig. 7. Graphical representation of principal component analysis for photosynthetic and bio-optical properties. Correlation circle in the plane by (A) the 2 first principal components, PC1 vs. PC2, and (C) the second and third principal components, PC3 vs. PC4. Projection of the stations composing the photosynthetic and bio-optical phytoplankton types (PBPT) for (B) PC1 vs. PC2 and (D) PC3 vs. PC4

mainly associated with bio-optical variables [Z_{eu} , $K_d(\text{PAR})$ and $a_d(440)$], while PC2 was mainly associated with physical variables (Sal , Z , SST and E_n). PC3 and PC4 explain only approximately 12% of the variation (Table 7). The projection of the different stations in relation to the PCs showed that some PBPT such as PBPT3 and PBPT7 were characterized by high values of $K_d(\text{PAR})$ and $a_d(440)$ and low values of Z_{eu} (PC1) (Fig. 9B,D). The rest of the PBPT did not show a clear association with a given PC.

Distribution of satellite phytoplankton size classes and field PBPT. The model of the distribution of phytoplankton types according to their size (Hirata et al. 2008) follows the satellite chl *a* distribution ob-

served in the mean images for the cruises (data not shown). Overlapping the PBPT distinguished here on this satellite classification, we can see that for a single phytoplankton size class, there are more than 1 PBPT and not always of the assumed size class; e.g. in the ‘micro class’ during spring, there are 3 PBPT (PBPT2, PBPT7 and PBPT8) which have a taxonomic composition that includes the nano size class as well. By contrast, PBPT3 was found prevalently in the ‘nano size’ but also in an area marked as ‘micro’ and in another marked as ‘pico’ (Fig. 10). Note, however, that our description of prevalent cell sizes for field data for the $<5 \mu\text{m}$ size fraction lacks information on the presence of cyanobacteria.

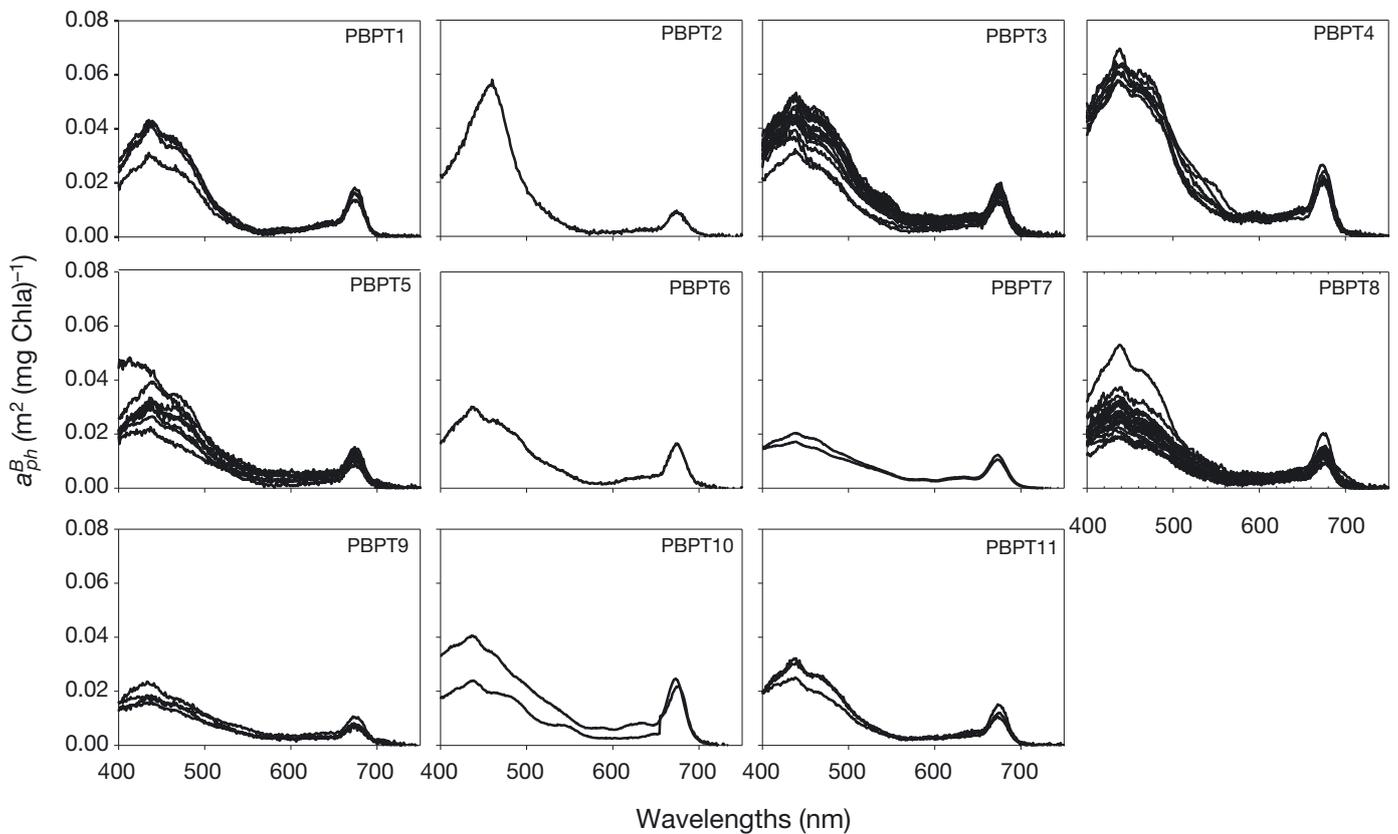


Fig. 8. Shapes of the specific absorption spectra of phytoplankton (normalized by chl *a*) of all the samples within each photosynthetic and bio-optical phytoplankton type (PBPT)

DISCUSSION

Oceanographic conditions in the study area

The large area of the Argentine shelf studied here—encompassing 18° of latitude (from temperate to subantarctic waters) and >1000 m difference in total depth (from the coast to the shelf break) during 3 seasons (spring, summer and winter)—showed a highly heterogeneous environment. From a biological point of view, the wide ranges in chl *a* concentration in these waters were caused by the occurrence of a large variety of different phytoplankton taxa. Among the physical and biological factors examined in this study, the most variable were the bio-optical properties (i.e. euphotic depth, attenuation coefficient and detritus absorption), while the main hydrographic measurements (e.g. *SST* and *Sal*) were more consistent (Table 7). The rich diversity in the phytoplankton communities found in the Argentine Sea could be attributed to the environmental heterogeneity, as shown in Table 3.

Chl *a* and primary production

The variability in the photosynthetic and bio-optical properties of phytoplankton produced large variations in the hourly and integrated primary production (Table 2). The $P_{Z,T}$ values reported here, considering the caveats of extrapolation of surface parameters, provide a first approximation of integrated production in different areas of the Argentine Sea. The positive and significant correlation found between surface *Chla* and p_0 for all stations in the 3 cruises ($r_s = 0.61$, $p < 0.05$) coincided with that in spring (Lutz et al. 2010). There were also significant correlations between *Chla* and $P_{Z,T}$ ($r_s = 0.59$, $p < 0.05$). Nevertheless, if one considers a linear direct relation between surface *Chla* and $P_{Z,T}$, the dispersion of the data is high; in some cases, the production is much lower (e.g. in GEF-1 Stn 35, in GEF-2 Stn 55 and in GEF-3 Stns 7 and 53) and in other cases is much higher (e.g. in GEF-1 Stns 8, 49, 71, 50 and 52) than that predicted from *Chla* (Fig. 3).

Table 6. Predominant taxonomic groups (dots) and range of percentage $Chla_{<5\mu m}$ within each photosynthetic and bio-optical phytoplankton type (PBPT). ND: no % $Chla_{<5\mu m}$ data were available

Size fraction		PBPT1	PBPT2	PBPT3	PBPT4	PBPT5	PBPT6	PBPT7	PBPT8	PBPT9	PBPT10	PBPT11
Range $Chla_{<5\mu m}$ (%)		74–100	26	45–100	16–100	41–94	100	ND	17–89	34–48	12–96	34–70
Bacillariophyceae	Ultra- <5 μm	•		•	•	•	•					
	Nano- 5–20 μm		•					•	•		•	•
	Micro- >20 μm		•			•				•		
Dinophyceae	Nano- 5–20 μm		•		•							•
	Micro- >20 μm		•						•			
Prymnesiophyceae <i>Emiliana huxleyi</i>	2–5 μm	•				•			•	•		
Other Prymnesiophyceae	2–5 μm			•	•	•						•
Chlorophyta (coccal)	~2 μm			•	•	•	•		•		•	

Phytoplankton types

The combination of the diverse taxa found and their physiological flexibility, according to our cluster analysis, revealed 11 PBPT which were associated with different photosynthetic and bio-optical properties (PCA, Fig. 7). This number of PBPT is higher than what is usually modeled by remote sensing according to cell size (pico-, nano- and microphytoplankton). Nevertheless, this grouping of the phytoplankton populations into 11 PBPT represents an important simplification with respect to the large amount of phytoplankton species present in the region; more than 330 species have been described for dinoflagellates alone (Balech 1988).

Table 7. Pearson correlation coefficients between main principal components (PC) and environmental properties. Percentage of variance accounted for by each component is shown in parentheses. Correlation coefficients were significant, above a critical value 0.30, with $p < 0.01$. Highest values of the coefficients in each component are shown in bold

Variable/ parameter	PC1 (29%)	PC2 (24%)	PC3 (12%)	PC4 (11%)
Z		0.70	-0.51	
Z_{eu}	0.92			
$K_d(\text{PAR})$	-0.86			0.36
MLD	0.56	-0.44		
SST		0.57	0.50	-0.36
Sal		0.77		0.38
E_n	-0.41	0.54		-0.36
$a_d(440)$	-0.65	-0.37		
Wi		0.33	-0.65	-0.49

Several studies have found a relationship between the bio-optical and physiological properties of phytoplankton and their cell size, e.g. Hoepffner & Sathyendranath (1992) showed that the highest values of $a_{ph}^B(\lambda)$ corresponded to small cells, and Bouman et al. (2005) showed that low values of P_m^B corresponded to microphytoplankton. Some clear cases of these relationships were also found here, e.g. PBPT1 formed mainly by small cells showed high specific absorption coefficients. Regarding ratios of absorption at different wavelengths ($a_{ph}(440)/a_{ph}(676)$ and $a_{ph}(490)/a_{ph}(555)$), PBPT3 (which had Prymnesiophyceae) had values close to those described for prymnesiophytes, and PBPT7 (composed of a bloom of diatoms) had values close to those reported for diatoms by Stuart et al. (2000). However, in other cases, variations in the photosynthetic and bio-optical properties did not always follow the expected trend according to the cell size. For example, PBPT5, with % $Chla_{<5\mu m}$ between 41 and 94%, was formed by a mixture of taxonomic groups, some of them ultraphytoplankton but others, microphytoplankton, and was characterized by high values of α^B and $a_{ph}^B(440)$ compared with other PBPT (Tables 4 & 7). This range of variability in bio-optical and photosynthetic parameters has been observed in other regions of the ocean. Claustre et al. (2005) found that for the North Atlantic (Iberian Peninsula and Azores archipelago), large cells had higher photosynthetic rates per unit carbon than expected; Cermeño et al. (2005) found that in the Ría de Vigo, large cells were more efficient at storing carbon than small groups; and Uitz et al. (2008) found that in the tropical North Atlantic and the equatorial and subequato-

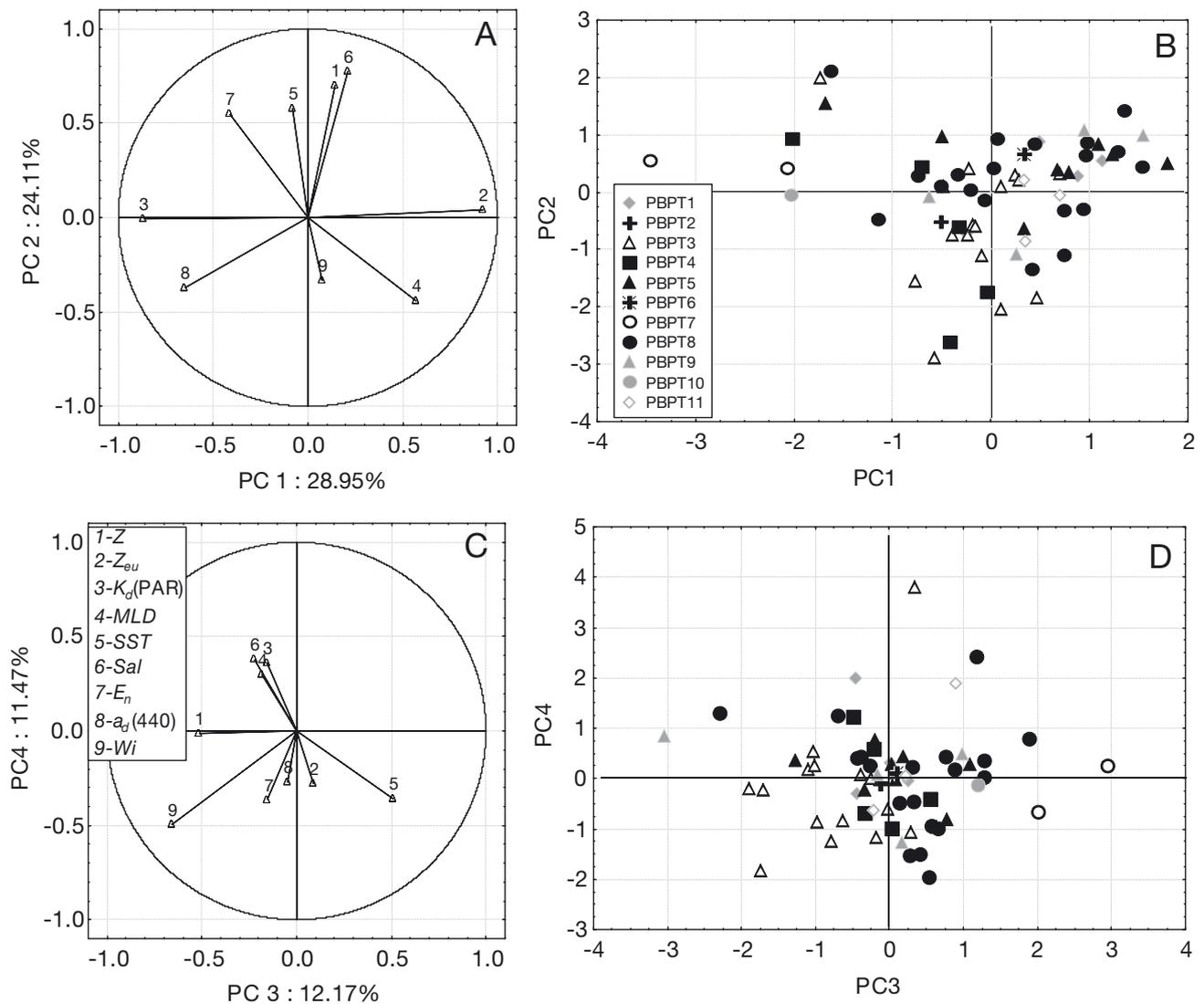


Fig. 9. Graphical representation of principal component analysis for environmental variables. Correlation circle in the plane by (A) the 2 first principal components, PC1 vs. PC2, and (C) PC3 vs. PC4. Projection of the stations composing the photosynthetic and bio-optical phytoplankton types (PBPT) for (B) PC1 vs. PC2 and (D) PC3 vs. PC4

rial Pacific, the microphytoplankton showed higher values of P_m^B than small-sized cells.

The high variation in photosynthetic parameters within the PBPT indicates that, at least for this data set, the relationship between bio-optical and photosynthetic characteristics was not strong. Therefore, the assignment of representative photosynthetic parameters for different phytoplankton types to model primary production in this region is quite difficult.

Some of the PBPT could be related to some particular environmental variables according to the PCA (e.g. PBPT7 and PBPT3 were associated with high values of $K_d(\text{PAR})$ and $a_d(440)$), although in most

cases, there was no clear pattern (Fig. 9). This would preclude the use of models of distribution of phytoplankton types according to physical characteristics of the environment, as has been proposed by Raitos et al. (2008).

Comparison between satellite and field phytoplankton types

Satellite and field distributions of phytoplankton types are difficult to compare with any accuracy since they are not defined by the same principles. In this case, we chose to distinguish our types using

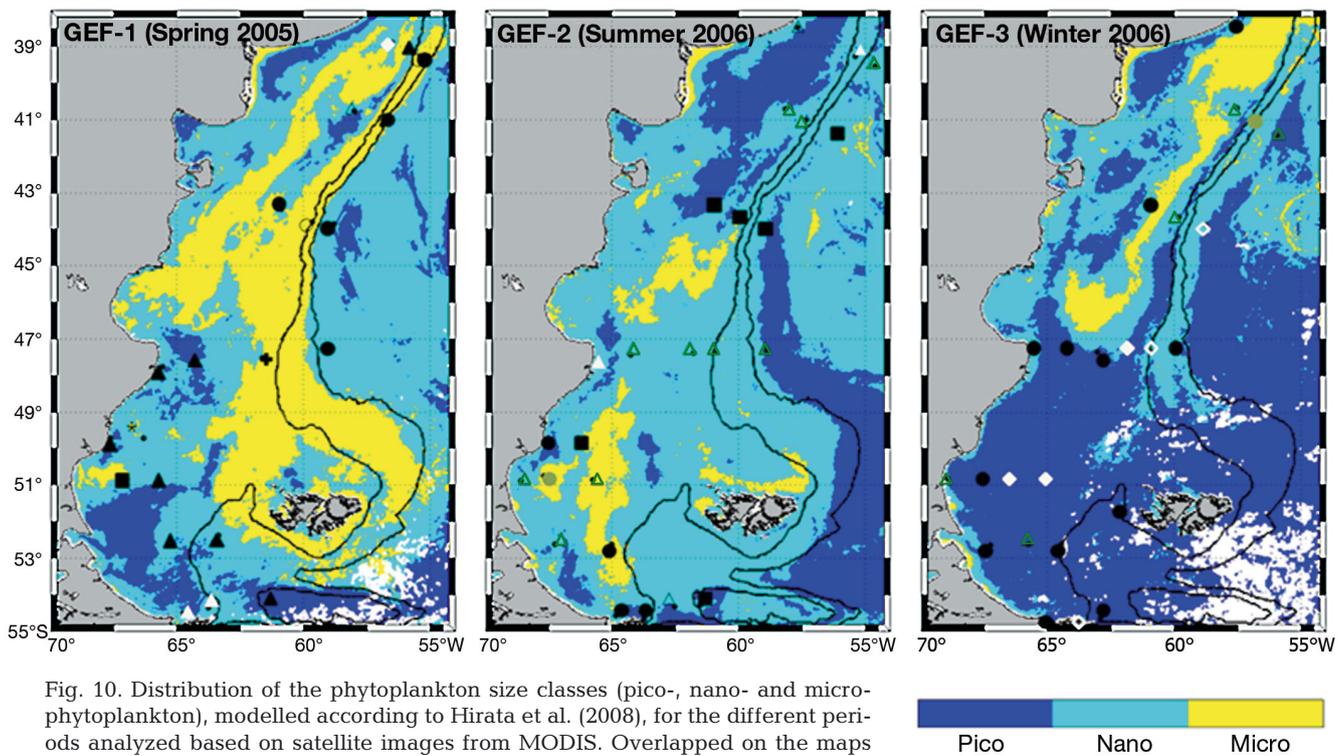


Fig. 10. Distribution of the phytoplankton size classes (pico-, nano- and micro-phytoplankton), modelled according to Hirata et al. (2008), for the different periods analyzed based on satellite images from MODIS. Overlapped on the maps are the distributions of the photosynthetic and bio-optical phytoplankton types distinguished in this study (see symbols in Fig. 6)

both photosynthetic and bio-optical characteristics. By way of comparison, the satellite model of Hirata et al. (2008) (one of the few that includes continental margins) is based only on phytoplankton cell size. Although there are some theoretical bases relating these properties (e.g. Aiken et al. 2007), these relationships do not seem to hold for the Argentine Sea because of the heterogeneity in environmental conditions and the wealth of phytoplankton taxonomic groups present. This complexity resulted in the distribution of more than 1 PBPT for a given size class (Fig. 10). A few field studies in well-characterized hydrographic sites have been able to represent the main groups of phytoplankton present using pigment composition alone (Carreto et al. 2003, Moreno et al. 2012). Nevertheless, we found that it is not easy to model the distribution of phytoplankton types using remote sensing in the whole Argentine Sea. This is due to the diversity of taxa present, from ultraphytoplankton (coccal, prymnesiophytes, diatoms) to microphytoplankton (dinoflagellates, diatoms), and their physiological flexibility, resulting in a high variability in bio-optical and photosynthetic properties, e.g. $a_{ph}^B(440)$ varied between 0.015 and 0.067 m^2 (mg chl a) $^{-1}$, and P_m^B varied between 0.68 and 10.05 mg C (mg chl a) $^{-1}$ h^{-1} .

Concluding remarks

The lack of a clear relationship between bio-optical variables and cell size, as found in this study, suggests that it is not always possible to model phytoplankton functional types using remote sensing based on considerations of cell size alone. Satellite models that identify phytoplankton types in this way (e.g. Alvain et al. 2005, Hirata et al. 2008) may not always be appropriate for highly heterogeneous regions such as the Argentine Sea.

In most of the PBPT, photosynthetic properties showed a higher degree of variability than bio-optical properties. Hence, it would be difficult to assign global photosynthetic parameters based on bio-optical characteristics derived from satellite models. Although we used data from extensive cruises during 3 seasons, this may nevertheless represent a somewhat limited number of data points to describe the variability found in the Argentine Sea. Additional field studies aimed at measuring photosynthetic rates should be performed in this zone of the southwestern Atlantic if we want to develop better biogeochemical models for the region that capture both the bio-optical as well as the physiological characteristics of this highly productive area.

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